

## Evaluation of the time of concentration estimation methods for small rural watersheds

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**Abstract:** The time of concentration (tc) is a widely used input parameter for the design of hydrographs, peak flow estimates, and runoff hydrological models. Conceptually, it is the average time it takes for water to flow from the most hydrologically distant location of the watershed to its outlet. The tc varies according to the characteristics of the basin, such as average slope, length and soil infiltration. It can be evaluated by calculating the time elapsed between the moment the effective rain stops and the end of the runoff identified by the hydrograph inflection point. When calculating the tc, there are some uncertainties that affect the performance of tc acquisition methods. The tc observations can generally be overestimated if they refer to runoff with a low hydraulic load. Because of that, the regular and effective assessment of watersheds is commonly used to validate the most appropriate tc estimation method. Among the equations used to estimate the tc of a watershed, this study aims to evaluate and compare the performance of twelve empirical and semi-empirical methods in three small rural watersheds in the state of São Paulo, Brazil. All watersheds had smaller areas than 81 km<sup>2</sup> and none of them are assessed. The tc estimates of the different models ranged between 1.08 h and 15.19 h, 1.00 h and 17.23 h, and 0.80 h and 8.20 h for the Barreiro, Copaiba and Lambari watersheds, respectively. The lowest tc value was estimated by the DNOS method and the longest one by the SCS Lag method for the Barreiro and Lambari watersheds. For Copaiba watershed, the lowest tc value was estimated by the DNOS method as well, but the highest tc value was derived by the Venturi equation. The tc estimates for the Lambari watershed were about twice higher than those of Copaiba watershed, although their areas were approximately similar. This difference among the estimates was strongly influenced by the explanatory variable of the thalweg models, followed by the mean slope. The results revealed that among the analysed formulas, the DNOS, Kirpich and Picking methods showed great similarity for presenting the smallest tc, causing the design of larger control structures. The SCS Lag and the Kinematic Wave method presented the highest tc, which can lead to lack of safety in hydraulic structures.

**Keywords:** *Rural watersheds, time of concentration, uncertainty, land use and land cover*

## 1. INTRODUCTION

For the design of hydraulic works and use of water resources it is necessary to determine what the available discharge is. In this sense, the effective discharge derivation depends on the proper estimation of the watershed time of concentration ( $t_c$ ). The  $t_c$  of a watershed can be defined as the time required for rainwater falling upon its most hydrologic distant region to runoff to its outlet (Tucci *et al.* 2009; Targa *et al.* 2012). In other words, it is the average time for water to reach the outlet as there is some dispersion throughout the watershed.

According to Araujo *et al.* (2011), factors such as basin area geometry, average terrain slope, sinuosity, soil infiltration, thalweg slope, among others, affect the time of concentration. Aiming to extend the knowledge about  $t_c$ , several studies have been carried out in different locations in which distinct empirical equations and methods have been used.

The determination of the  $t_c$  using empirical formulas is subject to the inaccuracies and uncertainties due to the type of flow that the formula seeks to represent. Consequently, continuous and appropriate hydrological monitoring is fundamental to validate the  $t_c$  estimation of watersheds (Almeida *et al.* 2014, Boulomytis *et al.* 2017).

Several researchers have developed empiric equations based on experimental and analytic methods to estimate the  $t_c$  (Kirpich 1940; Dooge 1956; Chow 1962). One of the most common empirical formulations in studies of this nature is the Kirpich method (Araujo *et al.* 2011). Although it is only applicable to very small watersheds (Kirpich 1940), in practice it is often used for watershed with a single main stream.

Because such numerous different equations, establishing a single and reliable estimation method for calculating the  $t_c$  becomes a challenge. The lack of possibility to perform direct measurements contributes to this scenario of multiple choices (Maia 2020).

In the current perspective, parameter options for calculating  $t_c$  can be estimated using more technologically advanced tools, such as the Geographic Information Systems (GIS). The GIS has useful tools to model information related to the hydrological factors of watersheds. It can be applied to subsidize the estimation of the surface runoff of watersheds and also calculation of hydrological groups, land use and land cover (LULC), previous soil moisture, and hydrological conditions that can be applied at the runoff Curve Number Method (Nasiri and Alipur 2014).

In this sense, this work aims to estimate the  $t_c$  of three small rural watersheds in the State of Sao Paulo using the ArcGIS software and twelve methods: Kinematic Wave, FAA, Kirpich, SCS Lag, Pasini, Ventura, DNOS, George Ribeiro, Ven te Chow, Johnstone, Corps of Engineers and Picking.

## 2. METHODOLOGY

### 2.1 Study area

We conducted the present work in three small rural watersheds (Barreiro, Lambari and Copaiba) within the Paraná basin located in the municipalities of Guzolandia, Casa Branca and Bastos, State of Sao Paulo, comprising areas of 80.35 km<sup>2</sup>, 66.98 km<sup>2</sup> and 65.30 km<sup>2</sup>, respectively (Figure 1).

### 2.2 Land use and land cover

The assessment of LULC information in these basins has been carried out based on MapBiomias (2020) regarding the data of 2019. We used the satellite images collected from LANDSAT 8 with spatial resolution of 30 m.

After the classification of MapBiomias imagery, we brought together some land use and land cover classes and renamed them due to their similar behaviour taking into consideration their hydrological features. For instance: the class “Farmland Formation” was grouped with the “Savannah Formation” class; the “Agriculture and Pasture” class was grouped with the “Pasture” class; and the class “Other non-vegetated area” was renamed to “Exposed Soil”.

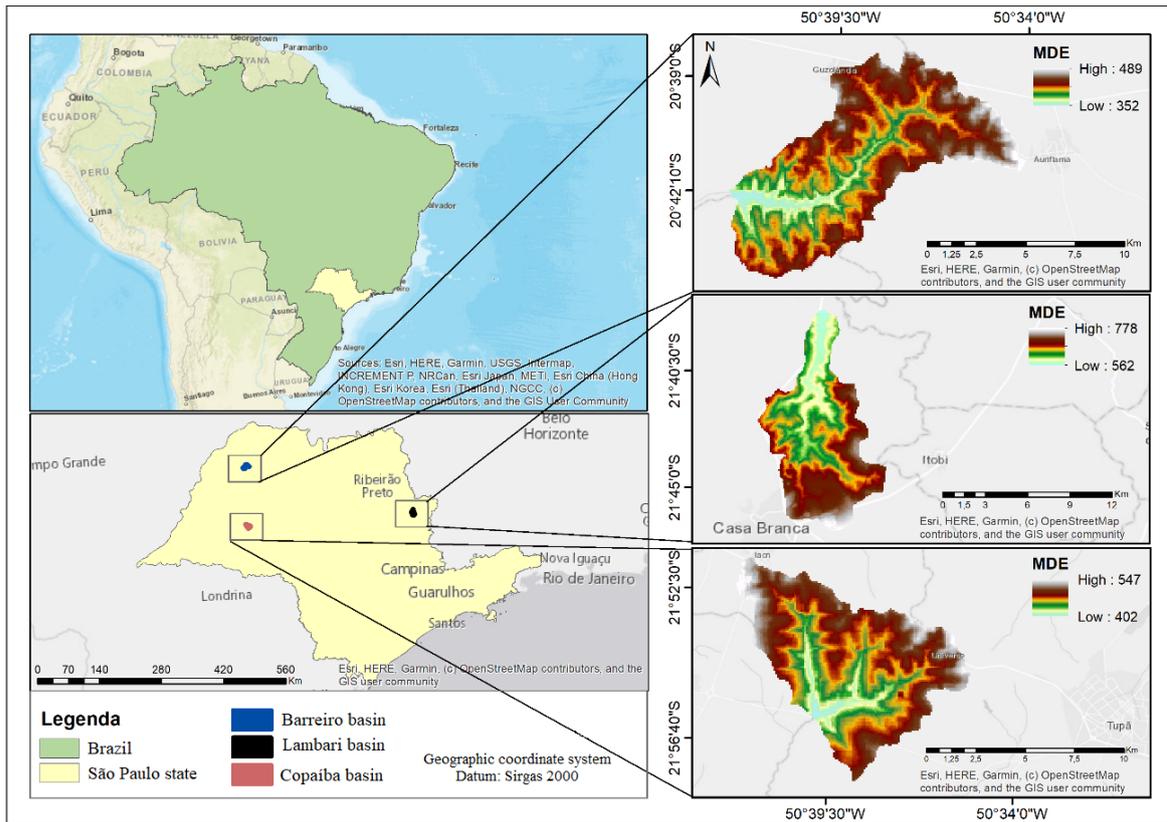


Figure 1. Spatial location of Barreiro, Lambari and Copaiba watersheds in the State of Sao Paulo, Brazil.

### 2.3 Time of concentration calculation

Twelve empirical and semi empirical methods were selected to calculate the times of concentration of the watersheds after an in-depth literature review. For the selection of methodologies, appropriate equations for rural watersheds were prioritized (Silveira 2005). To calculate the  $t_c$  (h), we used the following methods: Kinematic Wave, FAA, Kirpich, SCS Lag, Pasini, Ventura, DNOS, George Ribeiro, Vent Chow, Johnstone, Corps of Engineers and Picking (Table 1).

Table 1. Estimation methods used for the time of concentration.

Nome	Fórmula
Kinematic wave	$t_c = 7,35n^{0,6}i^{-0,4}L^{0,6}S^{-0,3}$
FAA	$t_c = 0,37(1,1 - C)L^{0,5}S^{-0,333}$
Kirpich	$t_c = 0,0663L^{0,77}S^{-0,385}$
SCS Lag	$t_c = 0,057(1000/CN - 9)^{0,7}L^{0,8}S^{-0,5}$
Ven te Chow	$t_c = 0,160L^{0,64}S^{-0,32}$
Johnstone	$t_c = 0,462L^{0,5}S^{-0,25}$
Corps of Engineers	$t_c = 0,191L^{0,76}S^{-0,19}$
Pasini	$t_c = 0,107A^{0,333}L^{0,333}S^{-0,5}$
Ventura	$t_c = 0,127A^{0,5}S^{-0,5}$
Picking	$t_c = 0,0883L^{0,667}S^{-0,333}$
DNOS	$t_c = 0,419k^{-1}A^{0,3}L^{0,2}S^{-0,4}$
George Ribeiro	$t_c = 0,222(1,05 - 0,2p)^{-1}LS^{-0,4}$

Source: Adapted from Silveira (2005).

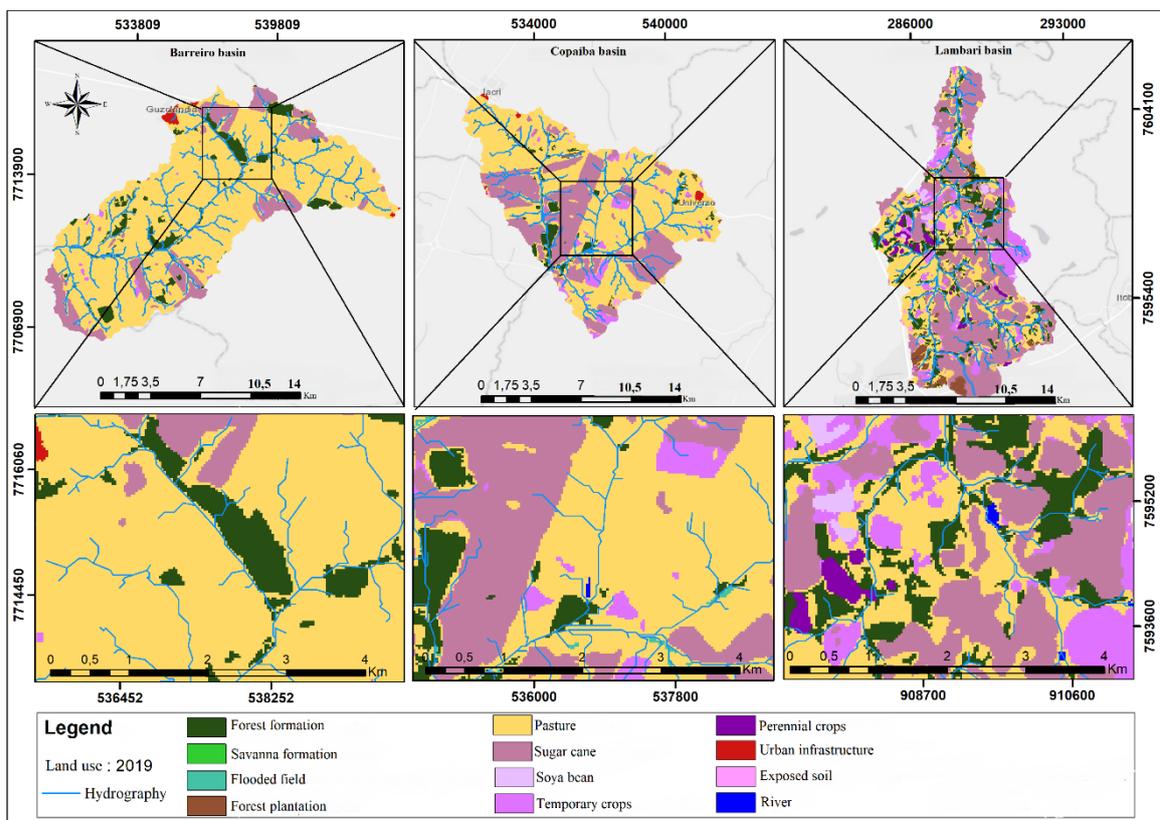
The parameters required for the  $t_c$  calculation were prepared in a GIS environment using the ArcGIS Desktop v.10.8 software. They were the area (A in km<sup>2</sup>), thalweg length (L in Kilometers) and slope (S in m/m). We adopted the rainfall intensity 35 mm/h according to McCuen et al. (1984). We obtained the values of the Manning's roughness coefficient, C parameter, the k factor of DNOS equation and the p factor of George

Ribeiro's equation according to Silveira (2005). The Curve Number (CN) was used by the SCS Lag equation. It was calculated for the three watersheds correlating the hydrological group of soil types and the LULC classes. The assignment of the CN values for each land cover class was based on the SCS standard table of values. After assigning values, the average CN values were achieved for each watershed according to its drainage area.

According to the Kinematic Wave, FAA, Kirpich, SCS Lag, Pasini, Ventura, DNOS and George Ribeiro equations, the length (L) must be determined for the source of the river and the slope must be calculated from the ratio between the maximum slope and length of course. On the other hand, Ven te Chow, Johnstone, Corps of Engineers and Picking mention L as the length of the main stream and S as its mean slope (Silveira 2005).

### 3. RESULTS AND DISCUSSION

Due to the LULC that represents the study watersheds, twelve classes were defined: forest formation, savanna formation, flooded field, forest plantation, pasture, sugarcane, soy beans, temporary crops, perennial crops, urban infrastructure, exposed soil and river. Figure 2 shows the LULC of the watersheds and Table 2 shows the watershed areas of each LULC using in 2019 data.



**Figure 2.** Land use and cover of the Barreiro, Copaiba and Lambari watersheds.

Pasture is the predominant LULC in the Barreiro and Copaiba watersheds, covering the rural areas of approximately 64.38 km<sup>2</sup> (80.12%) and 41.31 km<sup>2</sup> (63.26%), respectively. Sugarcane occupies the second largest area of 9.77 km<sup>2</sup> (12.16%) and 18.32 km<sup>2</sup> (28.06%) in the respective watersheds. On the other hand, Lambari watershed is mainly covered by sugarcane whose area is 23.80 km<sup>2</sup> (i.e. 35.53% of the total area). Pasture covers an area of 22.51 km<sup>2</sup> corresponding to 33.60% in this watershed.

The areas covered by natural vegetation (forest formation, savanna formation and flooded field) represent a smaller contribution to the landscape due to a secular deforestation, resulting in a large percentage of pasture and agricultural areas in the watersheds.

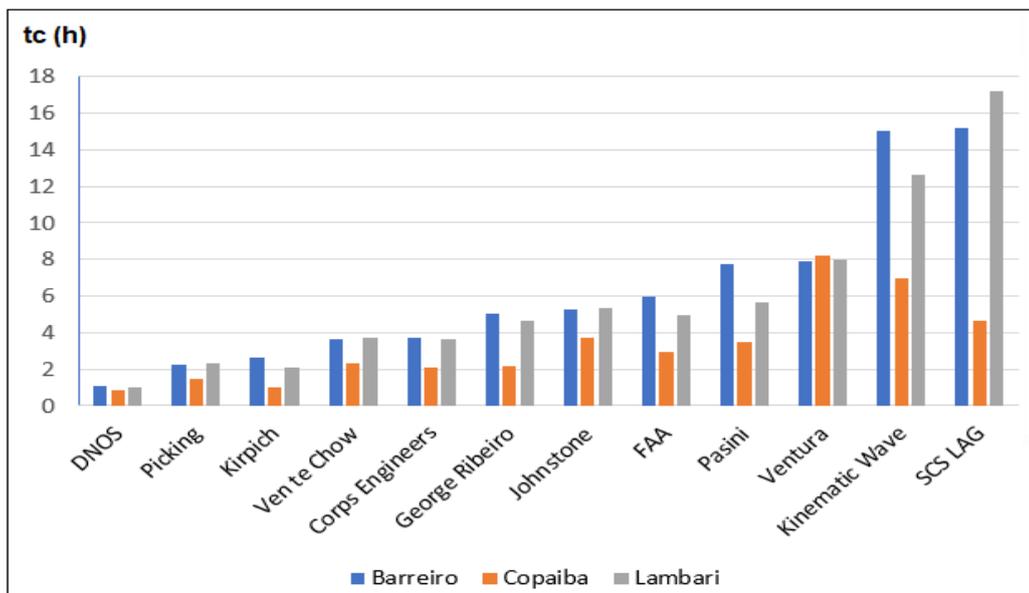
The watersheds are located on flat and gently undulating terrains consisting of two predominant types of tropical soil: red-yellow argisol (RYA) (in Copaiba and Lambari) and red latosol (RL) (in Barreiro). The RYA soils have been developed under the influence of crystalline rocks, which might have clay accumulation, and presence of iron oxides in subsurface horizons. The RL soils have deep developed profiles, a clayey texture and low natural fertility as their limiting factor. Nevertheless, they are suitable for agricultural and urban activities (Moraes *et al.* 1994; Pissarra *et al.* 2004).

**Table 2.** Percentage of basin areas classified into LULC.

LULC	Barreiro		Copaiba		Lambari	
	(km <sup>2</sup> )	(%)	(km <sup>2</sup> )	(%)	(km <sup>2</sup> )	(%)
Forest Formation	5.35	6.66	2.42	3.70	8.87	13.24
Savanna Formation	-	-	-	-	0.06	0.09
Flooded Field	-	-	0.28	0.43	-	-
Forest plantation	-	-	0.19	0.29	1.78	2.65
Pasture	64.38	80.12	41.31	63.26	22.51	33.60
Sugarcane	9.77	12.16	18.32	28.06	23.80	35.53
Soy beans	-	-	0.03	0.04	1.33	1.98
Temporary Crops	0.40	0.50	2.31	3.54	7.01	10.46
Perennial Crops	-	-	-	-	1,32	1.96
Urban Infrastructure	0.43	0.54	0.27	0.41	-	-
Exposed Soil	0.01	0.01	0.15	0.24	0.24	0.36
River	0.01	0.02	0.02	0.03	0.07	0.11
<b>Total</b>	<b>80.35</b>	<b>100</b>	<b>65.30</b>	<b>100</b>	<b>66.98</b>	<b>100</b>

The CN values properly represented the physical conditions of each watershed in terms of runoff, infiltration and LULC. The weighed CN values for Copaiba, Barreiro and Lambari watersheds were 64, 61 and 45, respectively. As expected, the highest CN values are related to the greater number of areas covered by pasture, corroborating to the fact that there is a greater runoff in Copaiba and Barreiro watersheds in relation to Lambari watershed. According to Raminhos (2002), rural watersheds have a higher infiltration capacity due to mainly vegetation cover, which is similar to a macro-roughness effect and thus, provides a higher resistance to runoff.

In Figure 3, we show the *t<sub>c</sub>* results, which were calculated by empirical methods. It should be observed that there is a great variability among the obtained values, and two extreme cases. The DNOS and Picking methods presented similar values, which can be explained by the similar parameters contained in their equations. On the other hand, the SCS Lag and Kinematic Wave methods presented the longest time of concentration for the three studied watersheds, which might be related to the fact that these methods incorporate constants on the features of land surface in their equations. These latest two methods have a semi-empirical deduction.



**Figure 3.** Time of concentration estimated for Barreiro, Lambari and Copaiba watersheds using 12 different methods.

The lowest *t<sub>c</sub>* value calculated for the Barreiro watershed was 1.08 h and the highest was 15.19 h (i.e. 14 times the lowest value). The mean value was 6.28 h and the median was 5.87 h, while the coefficient of variation

(CV) was 73%. The lowest  $t_c$  value of the Copaiba watershed was 1.00 h, while the highest was 17.23h, showing the increase of 17 times the lowest value. The mean was 5.93 h and the median was 4.81 h.

For Lambari watershed, the  $t_c$  value varied between 0.9 h and 8.2 h showing the increase of 10 times the lowest value, the mean of 3.32 h and the median was 2.62 h. The CV of the estimated  $t_c$  for the Copaiba and Lambari watersheds were 70 % and 79%, respectively.

For the three watersheds, the lowest  $t_c$  value was estimated by the DNOS method and the other two lower values were estimated by the Picking and Kirpich equations. In average, the second and third lowest  $t_c$  values were approximately twice the value calculated for the DNOS model for the Barreiro and Lambari watersheds. However, for the Copaiba watershed, the mean  $t_c$  value was 50% higher than the one achieved from the DNOS model.

The highest  $t_c$  for the Barreiro and Lambari watersheds were estimated by the SCS Lag model, while for the Copaiba watershed it was the  $t_c$  value calculated by the Ventura model.

Even though the areas of Copaiba (65.30 km<sup>2</sup>) and Lambari (66.98 km<sup>2</sup>) are approximately similar, the respective  $t_c$  values are totally distinct. It was verified that for the Lambari watershed the  $t_c$  value was almost twice the value of the ones found for the Copaiba watershed.

Regarding Copaiba watershed, the thalweg length (L) is 8.20 km, ratio between the maximum slope and L (M/L) is 0.06 m/m, and ratio between the maximum slope and the mean slope (M/S) is 0.0155 m/m. For Lambari watershed, L is 17.62 km, M/L is 0.04 m/m and M/S is 0.0171 m/m.

The  $t_c$  estimation is strongly influenced by L, and then, by the slope. For Copaiba and Lambari watersheds, the longest thalweg was determinant for the  $t_c$  estimation. Besides, the weighed CN value for Lambari watershed was 45, while for Copaiba it was 64. These values interfered in the  $t_c$  estimations using the SCS Lag method. Although the  $t_c$  was almost twice for Lambari watershed, comparing to the others, it was lower than the estimated  $t_c$  values by the kinematic wave and Ventura for the Copaiba watershed.

Araujo *et al.* (2011) analysed the  $t_c$  of watersheds and found that Ven Te Chow and Picking methods presented values which were very close to that obtained by the Kirpich method, even by varying the area and slope. According to Esteves (2003), Kirpich method presented a value close to the  $t_c$  observed in urban watersheds, as reported by Boulomytis *et al.* (2017). Kirpich method is the most widely used in simple structures in Brazil. However, the SCS Lag method overestimated  $t_c$ . Regarding the Kinematic Wave method, Tucci (1993) explains that it tends to overestimate the  $t_c$  by considering that runoff occurs on a plain watershed, which does not occur in practice where there are slopes. In this latest case, runoff get together after some meters in small ditches on the soil surface.

The water management authorities in Brazil generally use fully empirical equations such as DNOS, Picking and Kirpich, since they can provide greater reliability. Notwithstanding, these formulas can lead to the implementation of very large sized control structures, which can be is not advantageous for rural producers.

#### 4. CONCLUSION

The watershed sizes and runoff types are important information to define the most suitable methods for the  $t_c$  deviation. However, it is still not possible to recommend the most appropriate estimation method for watersheds without considering the observed hydrographs for comparison purposes.

The  $t_c$  estimation is essential for the implementation of control structures. In rural areas, it also helps the farmers to identify the most appropriate area for crops, based on the evaluation of the flood risk, which totally depends on a good estimation of the watershed  $t_c$ .

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#### REFERENCES

Almeida, I.K., Almeida, A.K., Anache, J. A. A., Steffen, J. L., and Sobrinho, T.A., 2014. Estimation on time of concentration of overland flow in watersheds: a review. *Geociências*, 33(4), 661-671.

- Araujo, B.A.M.D., Silveira, C.S., Souza, J.L., Maia Junior, J.V.F., Almeida, F.A.F., Studart, T.M.C., 2011. Análise do Tempo de Concentração em Função das Características Fisiográficas em Bacias Urbanas. In: XIX Simpósio Brasileiro de Recursos Hídricos, 2011, Fortaleza, 1-18.
- Boulomytis, V.T.G., Zuffo, A.C., Imteaz, M.A., Herrera, M.A.C., 2017. The effectiveness of the CN method in areas with saturated soil conditions. In: 22nd International Congress on Modelling and Simulation (MODSIM2017), 2017, Hobart, 1, 1663-1669.
- Chow, V.T., 1962. Hydrologic determination of waterway areas for the design of drainage structures in small drainage basins. Engineering Experiment Station Bulletin n.462. Urbana, In: University of Illinois College of Engineering, 104 p.
- Dooge, J.C.I., 1956. Synthetic unit hydrographs based on triangular inflow. Iowa State University, 103 p.
- Esteves, R.L. Mendiondo, E.M., 2003. Análise comparativa entre equações e observações do tempo de concentração em uma bacia urbana de Sao Carlos, SP. In: XV Simpósio Brasileiro de Recursos Hídricos, 2003, Curitiba.
- Kirpich, Z.P., 1940. Time of concentration in small agricultural watersheds, Civil Engineering, 10(6).
- Maia, J.C.P., 2020. Tempo de concentração em três bacias hidrográficas de Uberlândia, Minas Gerais. 48 f. Universidade Federal de Uberlândia, Uberlândia.
- Mapbiomas, 2020. Projeto MapBiomas – Coleção 5 da Série Anual de Mapas de Cobertura e Uso de Solo do Brasil. Available at <https://mapbiomas.org/colecoes-mapbiomas-br>. Accessed on 10 Apr. 2021.
- Mccuen, R.H., Wong, S.L., Rawls, W.J., 1984. Estimating urban time of concentration. Journal of Hydraulic Engineering, 110(7), 887-904.
- Moraes, I.C., Correa, E.A., Conceição, F.T., 1994. Análise da fragilidade ambiental utilizando técnicas em SIG: estudo de caso da bacia hidrográfica do córrego do Desemboque, Pirassununga –SP –Brasil. In: Simpósio Nacional de Geomorfologia, Brasília. Available at <http://lsie.unb.br/ugb/sinageo/8/10/9.pdf>. Accessed on 25 May 2021.
- Nasiri, A., Alipur, H., 2014. Determination the Curve Number catchment by using GIS and remote sensing. International Journal of Geological and Environmental Engineering, 8(5), 342–345.
- Pissarra, T. C. T., Politano, W., Ferraudo, A. S., 2004. Avaliação de características morfométricas na relação solo-superfície da bacia hidrográfica do córrego Rico, Jaboticabal (SP). Revista Brasileira de Ciência do Solo, 28.
- Raminhos, C, 2002. Experimentação e análise de resistência ao escoamento com macrorugosidades. Évora: Universidade de Évora. (Report).
- Silveira, A. L. L., 2005. Desempenho de fórmulas de tempo de concentração em bacias urbanas e rurais. RBRH: revista brasileira de recursos hídricos, 10(1), 5-23.
- Targa, M. D. S; Batista, G. T; Diniz, H. N; Dias, N. W; Matos, F. C., 2012. Urbanização e escoamento superficial na bacia hidrográfica do Igarapé Tucunduba, Belém, PA, Brasil. Revista Ambiente & Água, 7(2), 120-142.
- Tucci, C. E. M., 2009. Hidrologia: Ciência e Aplicação. 4 ed, Porto Alegre- RS:UFRG.